

Dimensional Analysis - Concepts

- **Physical quantities:** $R_j = v(R_j)[R_j] = \text{value} \cdot \text{unit}$, $j = 1, \dots, m$.
- **Units:** $[R_j] = F_1^{a_{1j}} \cdots F_n^{a_{nj}}$, F_1, \dots, F_n fundamental units.

- **Dimension matrix** of R_1, \dots, R_m : $A = \begin{bmatrix} a_{11} & \cdots & a_{1m} \\ \vdots & \ddots & \vdots \\ a_{n1} & \cdots & a_{nm} \end{bmatrix}$

- **Change of units** \Rightarrow change of values:

Lemma 1: $F_i = x_i \hat{F}_i$, $x_i > 0 \Rightarrow \hat{v}(R_j) = x_1^{a_{1j}} \cdots x_n^{a_{nj}} v(R_j)$

- **Dimensionless combination:** $\pi = R_1^{\lambda_1} \cdots R_m^{\lambda_m}$ if $\vec{\lambda} \neq 0$, $[\pi] = 1$
- **Dimensionally independent** R_1, \dots, R_s if no dimensionless comb'ns exist
- **Physical relations** $\Phi(R_1, \dots, R_m) = 0$ are **dimensionally consistent**, i.e.

$$\Phi(v(R_1), \dots, v(R_m)) = 0 \Leftrightarrow \Phi(\hat{v}(R_1), \dots, \hat{v}(R_m)) = 0$$

for all changes of units \hat{F}_i . (consistent under change of units)

Dimensional Analysis - Buckingham's pi-theorem

- (A1) F_1, \dots, F_n are fundamental units
- (A2) R_1, \dots, R_m are physical quantities
- (A3) $\Phi(R_1, \dots, R_m) = 0$ is dimensionally consistent.

Lemma 2: Let $r = \text{rank } A$, then R_1, \dots, R_m have $m - r$ independent dimensionless combinations.

OBS: The rank = number of linearly independent columns in the matrix.

Buckingham's pi-theorem:

If (A1) – (A3) hold, then there are $m - r$ independent dimensionless combinations, and for any set of $m - r$ independent dimensionless combinations π_1, \dots, π_{m-r} , there is a relation Ψ such that

$$\Phi(R_1, \dots, R_m) = 0 \quad \Leftrightarrow \quad \Psi(\pi_1, \dots, \pi_{m-r}) = 0,$$

where $r = \text{rank } A$ and A is the $n \times m$ dimension matrix of R_1, \dots, R_m .

It remains to prove **Lemma 2** and the **Pi-theorem**.

Scaling and non-dimensionalizing

Produce dimensionless $O(1)$ variables and dim.less eq'ns with terms $\lesssim 1$

Scaling a variable u^* : $u^* = Uu$ where

scaled variable: $u \sim 1$, $[u] = 1$

scaling constant: $U \sim \max |u^*|$, $[U] = [u^*]$

Scaling/nondimensionalizing an equation:

- 1 scaling all variables in the equation
- 2 dividing the resulting equation by \sim biggest coefficient.

Finding scales:

- look for combinations of the parameters
- balance 2 dominating ("biggest") terms in the equation (using that all scaled variables should be $O(1)$)
- solve a reduced problem to find estimates

- typical time scale for $u^*(t^*)$:
$$T = \frac{\max |u^*|}{\max \left| \frac{du^*}{dt^*} \right|}$$

Remarks on scaling

Remark 1: 2 dominating terms balanced

⇒ their coefficients are equal and \sim biggest in equation.

⇒ dividing scaled equation by this coefficient:

All variables and coefficients become
dimensionless,

$\lesssim 1$, and

2 coefficients = 1.

Remark 2: Different situations ⇒ different scales

$\max |x^*|$, $\max |t^*|$, $\max |u^*|$, etc., and the dominating terms in the equation depend on the situation.

Remark 3: Advantages of scaling:

- minimize the number of parameters/coefficients (experiments!!),
- normalize all variables and coefficients,
- reduce round-off errors in subsequent numerical calculations,
- make small terms visible ⇒ easy to do approximations/perturbation.

Regular Perturbation

Given scaled(!) equation: $\ddot{x} = -\frac{1}{(1 + \varepsilon x)^2}$, $0 < \varepsilon \ll 1$.

1. **Perturbation Assumption:** $x = x_0 + \varepsilon x_1 + \varepsilon^2 x_2 + \dots$
2. Insert into equation, expand as **power series in ε** :

$$\begin{aligned}\ddot{x}_0 + \varepsilon \ddot{x}_1 + \varepsilon^2 \ddot{x}_2 + \dots &= -\frac{1}{\left(1 + \varepsilon(x_0 + \varepsilon x_1 + \dots)\right)^2} \\ &= -1 + 2\varepsilon(x_0 + \varepsilon x_1 + \dots) - 3\varepsilon^2(x_0 + \varepsilon x_1 + \dots)^2 + \dots \\ &= -1 + \varepsilon 2x_0 + \varepsilon^2(2x_1 - 3x_0^2) + \dots\end{aligned}$$

3. Equate terms of **same order in ε** \rightarrow equations for x_0, x_1, \dots :

$$O(1): \quad \ddot{x}_0 = -1$$

$$O(\varepsilon): \quad \ddot{x}_1 = 2x_0$$

$$O(\varepsilon^2): \quad \ddot{x}_2 = 2x_1 - 3x_0^2$$

5. Solve these equations **recursively** for x_0, x_1, x_2, \dots .

Singular Perturbation

Signs:

- Multiple time/space scales
- Initial/boundary layers
- Small parameter multiplying principal term
- Naive approximation changes problem completely

Facts:

- No single scale is good for complete resolution of problem
- Different regions, different scales, different (re)scaled equations
- Scales found by balancing terms in equation
- In each region regular perturbation works
- Matching conditions between solutions of different regions

Singular Perturbation – first approximation

$$\varepsilon y'' + 2y' + y = 0, \quad 0 < x < 1; \quad y(0) = 0, \quad y(1) = 1; \quad 0 < \varepsilon \ll 1.$$

1. **Guess** where boundary layer is: $x = a$. Here $a = 0$.
2. **Outer solution** y_O . Set $\varepsilon = 0$ and solve **equation** and boundary condition **outside** boundary layer:

$$2y'_O + y_O = 0; \quad y_O(1) = 1 \implies y_O(x) = e^{\frac{1}{2} - \frac{x}{2}}.$$

3. Find **length of boundary layer** δ (the other *consistent* space scale) by **balancing terms** $\rightarrow \dots \delta = \varepsilon$.
4. **Rescale** equation: $(x, y) = (\delta\xi, Y) \rightarrow Y''(\xi) + 2Y'(\xi) + \varepsilon Y(\xi) = 0$
5. **Inner solution** y_I . Set $\varepsilon = 0$ and solve **rescaled equation** and boundary condition **inside** boundary layer.

$$y''_I + 2y'_I = 0, \quad y_I(0) = 0 \implies y_I(\xi) = C(1 - e^{-2\xi}).$$

6. **Matching**. $y_O \approx y_I$ in **intermediate region** $\xrightarrow{\varepsilon \rightarrow 0} \dots C = e^{\frac{1}{2}}$ (approx'n!)
7. **Uniform solution**: $y_U(x) = y_O(x) + y_I\left(\frac{x}{\delta}\right) - \lim_{x \rightarrow 0} y_O(x)$

Equilibrium points

1. **Equilibrium point** = constant solution u_e (e.g. of ODEs or PDEs)
2. An equilibrium point u_e is **stable** if all solutions starting near u_e , remain near u_e for all $t \geq 0$.
3. **Linear stability analysis**
 - 1 Write solution $u = u_e + \tilde{u}$, \tilde{u} small perturbation
 - 2 Linearize equation(s) about u_e :
insert $u = u_e + \tilde{u}$ into equation(s)
drop small(=non-linear in \tilde{u}) terms
Result: linear equation(s) for \tilde{u} ,
with equilibrium point $\tilde{u}_e = 0$.
 - 3 Check stability of $\tilde{u}_e = 0$ (linearized equation(s)!!)
 - 4 Conclusion: $\tilde{u}_e = 0$ stable/unstable indicate that u_e stable/unstable.
4. Over time all physical systems tend to be at their stable equilibrium solutions! (... always small disturbances ...)

Aggregation of Amoeba

Background: *Lack of food* \rightarrow amoeba produce attractant and aggregate.

Question:

Can onset of aggregation be caused by simple, unintelligent mechanism?

Model near onset of aggregation:

- **Physical quantities:**

$a(x, t)$, $c(x, t)$ = amoeba, attractant densities; parameters

- **Modelling (conservation+diffusion+attraction+production+decay):**

$$(1) \quad a_t = \frac{\partial}{\partial x} (ka_x - lac_x), \quad c_t = Dc_{xx} + q_1a - q_2c.$$

- **Equilibrium points** (=constant solutions):

Constants (a_0, c_0) such that $q_1a_0 = q_2c_0$.

- **Linearize equation around (a_0, c_0) :**

$a = a_0 + \tilde{a}$, $c = c_0 + \tilde{c}$; \tilde{a}, \tilde{c} small; drop small terms

$$(2) \quad \tilde{a}_t = \frac{\partial}{\partial x} (k\tilde{a}_x - la_0\tilde{c}_x), \quad \tilde{c}_t = D\tilde{c}_{xx} + q_1\tilde{a} - q_2\tilde{c}.$$

Aggregation of Amoeba

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- **Particular solutions of (2):** Fourier modes/eigenfunctions

$$(\tilde{a}, \tilde{c}) = e^{\alpha t} \cos(\beta x) (C_1, C_2)$$

solve (2) iff

- $\alpha^2 + b\alpha + c = 0$ for

$$b = k\beta^2 + D\beta^2 + q_2 \quad \text{and} \quad c = kq_2\beta^2 + kD\beta^4 - q_1la_0\beta^2,$$

- (C_1, C_2) satisfy two linear equations (last time).

Every $\beta \in \mathbb{R} \rightarrow$ two real α + linear subspace of solutions (C_1, C_2)

- **Boundedness of solutions of (2):**

$$(\tilde{a}, \tilde{c}) \text{ bounded} \Leftrightarrow \alpha \leq 0 \Leftrightarrow c \geq 0 \Leftrightarrow \boxed{kD\beta^2 + kq_2 \geq q_1la_0}$$

- **Unboundedness:** $kq_2 < q_1la_0 \Rightarrow (\tilde{a}, \tilde{c})$ blow up at ∞ for $\beta \ll 1$
 - Take C_1, C_2 arbitrarily small $\Rightarrow (0, 0)$ unstable equilibrium pt. of (2)
 - Indicate that $(a_0, c_0) (\Leftrightarrow (\tilde{a}, \tilde{c}) = (0, 0))$ **unstable equi. pt.** of (1)

Conservation in Continuum Mechanics

The transport theorem: $R(t) = \{x(t) : \dot{x} = \vec{v}(x, t), x(t_0) = x_0 \in R(t_0)\}$

$$\frac{d}{dt} \int_{R(t)} f(x, t) dx \Big|_{t=t_0} = \frac{d}{dt} \int_{R(t_0)} f(x, t) dx \Big|_{t=t_0} + \int_{\partial R(t_0)} f(x, t_0) (\vec{v} \cdot \vec{n}) d\sigma$$

Conservation of mass and momentum:

$$(1) \quad \frac{d}{dt} \int_R \rho dx + \int_{\partial R} \rho (\vec{v} \cdot \vec{n}) d\sigma = \int_R q dx.$$

$$(2) \quad \frac{d}{dt} \int_R \rho \vec{v} dx + \int_{\partial R} \rho \vec{v} (\vec{v} \cdot \vec{n}) d\sigma$$

Transp. thm. $\frac{d}{dt} \int_{R(t)} \rho \vec{v} dx$ Newton's 2nd law $\int_{R(t)} \vec{f}_B dx$ + $\int_{\partial R(t)} \vec{f}_S d\sigma$
body forces surface forces

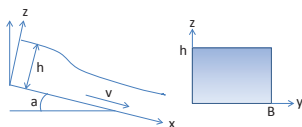
Newtonian fluid: $\vec{f}_S = T \cdot \vec{n}$, $T_{ij} = -\left(\rho + \frac{2}{3}\mu \nabla \cdot \vec{v}\right)\delta_{ij} + \mu\left(\frac{\partial v_i}{\partial x_j} - \frac{\partial v_j}{\partial x_i}\right)$

Differential form - the Navier-Stokes equations:

$$(1') \quad \rho_t + \nabla \cdot (\rho \vec{v}) = q$$

$$(2') \quad (\rho v_i)_t + \nabla \cdot (\rho v_i \vec{v}) = f_{B,i} - \frac{\partial p}{\partial x_i} + \mu\left(\nabla^2 v_i + \frac{1}{3} \frac{\partial}{\partial x_i} (\nabla \cdot \vec{v})\right), \quad i = 1, 2, 3$$

Flow in rivers



Assumptions:

① v, h depend only on x, t ; a small; $\rho = \text{const}$.

② Dominant forces in x -direction:

$$\text{Gravity } \vec{f}_g \cdot \vec{e}_x = \rho g \sin a$$

$$\text{Hydrostatic pressure } \vec{f}_p \cdot \vec{e}_x \approx -\rho g (h - z) (\vec{n} \cdot \vec{e}_x)$$

$$\text{Bottom friction } \vec{f}_f \cdot \vec{e}_x = -\rho C_f v^2$$

Control volume:

$$R = \{(x, y, z) : x \in [x_0, x_0 + \Delta x], y \in [0, B], z \in [0, h(x, t_0)]\}$$

Conservation in R of mass and momentum in x -direction:

$$\frac{d}{dt} \int_R \rho dx + \int_{\partial R} \rho (\vec{v} \cdot \vec{n}) d\sigma = 0,$$

$$\frac{d}{dt} \int_R \rho v dx + \int_{\partial R} \rho v (\vec{v} \cdot \vec{n}) d\sigma = \int_R \vec{f}_g \cdot \vec{e}_x dx + \int_{\partial R} (\vec{f}_p + \vec{f}_f) \cdot \vec{e}_x d\sigma.$$

Flow in rivers

Compute all integrals, divide by common factor ρB :

$$\frac{d}{dt} \int_{x_0}^{x_0 + \Delta x} h \, dx + [(vh)(x_0 + \Delta x) - (vh)(x_0)] = 0,$$

$$\begin{aligned} & \frac{d}{dt} \int_{x_0}^{x_0 + \Delta x} vh \, dx + [(v^2h)(x_0 + \Delta x) - (v^2h)(x_0)] \\ &= \int_{x_0}^{x_0 + \Delta x} gh \sin a \, dx - \frac{g}{2} [h^2(x_0 + \Delta x) - h^2(x_0)] - \int_{x_0}^{x_0 + \Delta x} C_f v^2 \, dx. \end{aligned}$$

Divide by Δx , let $\Delta x \rightarrow 0$:

$$h_t + (vh)_x = 0,$$

$$(vh)_t + \left(v^2h + \frac{g}{2}h^2\right)_x = gh \sin a - C_f v^2.$$

This is the shallow water equations or St. Venant system.